

## **The relationship of multispectral satellite imagery to immediate fire effects**

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Andrew T. Hudak<sup>1\*</sup>, Penelope Morgan<sup>2</sup>, Michael J. Bobbitt<sup>2</sup>, Alistair M.S. Smith<sup>2</sup>, Sarah A. Lewis<sup>1</sup>, Leigh B. Lentile<sup>2</sup>, Peter R. Robichaud<sup>1</sup>, Jess T. Clark<sup>3</sup>, and Randy A. McKinley<sup>4</sup>

<sup>1</sup>Forest Service, U.S. Department of Agriculture, Rocky Mountain Research Station, Moscow, Idaho

<sup>2</sup>University of Idaho, Department of Forest Resources, Moscow, Idaho

<sup>3</sup>RedCastle Resources, Inc. contractor to the Forest Service, U.S. Department of Agriculture, Remote Sensing Applications Center, Salt Lake City, Utah

<sup>4</sup>Scientific Applications International Corporation contractor to the U.S. Geological Survey, Center for Earth Resources Observation and Science, Sioux Falls, South Dakota

### **Abstract**

The U.S. Forest Service Remote Sensing Applications Center (RSAC) and the U.S. Geological Survey Earth Resources Observation and Science (EROS) Data Center produce Burned Area Reflectance Classification (BARC) maps for use by Burned Area Emergency Response (BAER) teams in rapid response to wildfires. BAER teams desire maps indicative fire effects on soils, but photosynthetic, non-photosynthetic vegetation, and other materials also affect the spectral properties of post-fire imagery. Our objective was to assess burn severity both remotely and on the ground at six 2003 and two 2004 wildfires. We measured or calculated fire effects variables at 418 plots, nested in 50 field sites, located across the full range of burn severities observed at the 2003 Black Mountain, Cooney Ridge, Robert, and Wedge Canyon wildfires in western Montana, the 2003 Old and Simi wildfires in southern California, and the 2004 Porcupine and Chicken wildfires in interior Alaska. We generated the Normalized Burn Ratio (NBR), delta Normalized Burn Ratio (dNBR), Relative dNBR (RdNBR), Normalized Difference Vegetation Index (NDVI), and delta NDVI (dNDVI) burn severity indices from Landsat 5 Thematic Mapper (TM) imagery across these eight wildfires. The NBR correlated best with the fire effects measures but insignificantly, meaning other indices could act as suitable substitutes. The overstory (trees in Montana and Alaska, shrubs in California) measures appear to have more influence on the image variables, followed by understory and surface cover measures. Exposed mineral soil and soil water repellency were poorly correlated with the image variables, while green vegetation was most highly correlated. The BARC maps are more indicative of post-fire vegetation condition than soil condition. Our future research will focus on spectral mixture analysis (SMA) because it acknowledges that pixel reflectance is fundamentally a mixture of charred, dead, green, and nonphotosynthetic vegetation, soil, rock and ash materials that are highly variable at fine scales.

**Keywords:** burn severity, char, Landsat, remote sensing, soils, spectral mixture analysis, vegetation, wildfire

\*Corresponding author contact information: [ahudak@fs.fed.us](mailto:ahudak@fs.fed.us), 208-883-2327 (phone), 208-883-2318 (fax)

## Introduction

Large wildfires have occurred and will continue to occur often in ecosystems of the United States, especially the drier ecosystems of the western USA (Morgan *et al.* 2003). Wildfires are an essential component of these ecosystems, but have become increasingly expensive to suppress as human development expands the wildland-urban interface (WUI), where local inhabitants would rather protect their property than maintain long-term ecological processes. The U.S. Forest Service and other government land management agencies annually expend millions of dollars to suppress wildfires that endanger people or their property. Additional millions are spent by Burned Area Emergency Response (BAER) teams to rehabilitate recently burned areas, especially severely burned areas that are especially vulnerable to erosion, sedimentation of water supplies, or encroachment by undesirable invasive species. Efforts to increase the efficiency of post-fire rehabilitation treatments rely in large part on accurate maps of where severely burned areas occur across the landscape, along with their appropriate use in their assessments.

Increasingly, BAER team leaders turn to Burned Area Reflectance Classification (BARC) maps for planning emergency response to wildfires. BARC maps are produced as rapidly as possible by the Remote Sensing Applications Center (RSAC) on U.S. Forest Service (USFS) managed lands, and U.S. Geological Survey (USGS) Center for Earth Resources Observation and Science (EROS) on Department of Interior managed lands. BARC maps are so named to distinguish themselves from burn severity maps, which require ground validation by the BAER teams to determine appropriate thresholds for distinguishing severely burned areas from areas only moderately or lightly burned. BARC products are derived from either the Normalized Burn Ratio (NBR) or delta NBR (dNBR) indices (Key and Benson 2003b, van Wagtenonk *et al.* 2004). Although dNBR is the default choice as a burn severity indicator, Bobbe *et al.* (2003), in a field validation of BARC products, found dNBR to be no more accurate than NBR. They called for further assessments across a greater range of ecosystems, a need that prompted this study.

Imagery from low spatial resolution satellite sensors such as MODIS, SPOT-VEG, and AVHRR have been widely used for regional measures of burned area (Barbosa *et al.* 1999, Stroppiana *et al.* 2002). Landsat imagery is the default choice for mapping smaller fires key to local management needs (Clark *et al.* 2003), but satellite sensors such as SPOT, ASTER, and other imagery are important supplements. Regional burn area products (e.g., GBA2000, GLOBSCAR, etc.) have been heavily evaluated (Gregoire *et al.* 2003), and there currently exists extensive literature that has investigated the application of remote sensing techniques to measure the area burned using a wide variety of satellite sensors at both moderate (Smith *et al.* 2002, Hudak and Brockett 2004) and broad scales (Barbosa *et al.* 1999, Zhang *et al.* 2003).

In contrast, the remote assessment of the fire severity from aerial and satellite imagery remains relatively under researched (Lentile *et al.* 2006). Indeed, the relationships between spectral properties of burned areas and field measures of burn severity have been evaluated in few studies (Landmann 2003, van Wagtenonk *et al.* 2004, Smith *et al.* 2005, Lewis *et al.* 2006, Robichaud *et al.* In Press). The remote assessment of burn severity is expected to be highly dependent on the acquisition date of the image following the fire, as spectral evidence such as soil charring the quantity and character of the ash produced by the fire will be quickly altered by meteorological processes and vegetation regrowth (Robinson 1991). Therefore, the utility of NBR, dNBR, and

other indices with potential for mapping burn severity needs to be quantitatively tested across a variety of image types and acquisition dates. Preferably, maps of specific fire effects (e.g., cover of exposed soil, organic matter, or vegetation) with biophysical relevance might be produced, instead of maps keyed using loose burn severity definitions that are less interpretable and useful (Lentile *et al.* 2006), and may even run the risk of misinterpretation or misuse. To obtain this information, we conducted an extensive field validation project at four wildfires in western Montana, two in southern California, and two in interior Alaska. Our objective was to assess how well burn severity indices derived from Landsat Thematic Mapper (TM) imagery relate to a suite of immediate post-fire effects measured on the ground.

## **Methods**

### **Wildfires Sampled**

We characterized fire effects across the full range of burn severity observed in the field, as soon as possible after eight large wildfire events in 2003 and 2004 (Fig. 1). The Black Mountain and Cooney Ridge wildfires, located west and east of Missoula, Montana, respectively, together burned over 10,000 ha during much of August and into September, 2003. Beginning in mid-July and for the next two months, the Robert and Wedge Canyon wildfires west of Glacier National Park burned nearly 45,000 ha combined. In less than two weeks between late October and early November, 2003, the Old and Simi wildfires north of San Bernadino and Simi Valley, respectively, collectively burned over 80,000 ha in southern California. The largest 2004 wildfires in the United States were in interior Alaska, the largest of which was the Taylor Complex southeast of Fairbanks. The Porcupine and Chicken wildfires sampled in this project in late July expanded along with the Billy Creek and Gardiner Creek fires to the west to eventually merge into the >520,000 ha Taylor Complex portrayed in Fig. 1.

Before entering an active fire zone and upon leaving, we followed established safety protocols and communicated our location and intention with the active Incident Command. The importance of strict adherence to wildfire safety procedures cannot be overemphasized (Lentile *et al.* *In Press*). One advantage of our rapid response research project was our ability to obtain fire progression records from the Incident Command GIS team, while on these active incidents. From these records, or by asking fire personnel who were on the ground at the time of burning, we established the burn dates for all of the plots we sampled. This later allowed us to calculate how soon after each plot burned were the fire effects characterized.

(Figure 1 about here)

### **Image Processing**

The spectral channels, spatial resolution and coverage, temporal frequency, and low cost of Landsat TM make it the preferred image sensor for rapid production of BARC maps by RSAC and EROS. Satellite images are selected based on the availability of a cloud-free scene as soon as possible after the need for BARC products is identified.

All imagery used in this paper were from the Landsat 5 TM sensor (Table 1), eliminating sensor type as a source of variation in the analysis. All the images were provided by either RSAC (Montana and California fires) or EROS (Alaska fires). Each was already rectified geometrically and radiometrically, and calibrated to top-of-atmosphere reflectance, following accepted preprocessing procedures ([http://landcover.usgs.gov/pdf/image\\_preprocessing.pdf](http://landcover.usgs.gov/pdf/image_preprocessing.pdf)).

(Table 1 about here)

We calculated NBR (Eq. 1) because it is the most applied burn severity index, together with its corresponding delta index, delta NBR (dNBR; Eq. 2), developed by Key and Benson (2003b). We also assessed the Relative dNBR (RdNBR; Eq. 3), which transforms the dNBR to a relative scale and can remove heteroscedasticity from the distribution of dNBR values, which improved burn severity classification accuracy in the Sierra Nevada mountains of California (Miller and Thode, *In Press*). In addition, we included the NDVI (Eq. 4) because of its broad use in remote sensing across most satellite sensors, and the corresponding change index, delta NDVI (dNDVI; Eq. 5).

$$NBR = (NIR - SWIR) / (NIR + SWIR) \quad (1)$$

$$dNBR = NBR_{prefire} - NBR_{postfire} \quad (2)$$

$$RdNBR = (NBR_{prefire} - NBR_{postfire}) / \sqrt{(|NBR_{prefire}| / 1000)} \quad (3)$$

$$NDVI = (NIR - RED) / (NIR + RED) \quad (4)$$

$$dNDVI = NDVI_{prefire} - NDVI_{postfire} \quad (5)$$

In the above formulas, RED denotes the red band, Landsat band 3; NIR denotes the near infrared band, Landsat band 4; and SWIR denotes the short wave infrared band, Landsat band 7. These burn severity indices were created in ERDAS Imagine.

We also applied spectral mixture analysis (SMA) to the six reflectance bands of each post-fire Landsat TM image to estimate green vegetation (green), nonphotosynthetic vegetation (NPV) (brown), and char (black) fractional cover, using the linear SMA tool in the ENVI software package (RSI, Boulder, CO). SMA is an established remote sensing method that has been applied to both delineate areas burned and calculate the fractional cover of vegetation and char within burned pixels (Wessman *et al.* 1997, Cochrane and Souza 1998, Smith *et al.* *In Press*). In SMA, the solution of a linear model enables the calculation of the relative proportions that a given cover type contributes to a pixel's reflectance. The linear mixture model is defined according to Eq. 6 by (Drake *et al.* 1999):

$$R_i = \sum_{c=1}^n (r_j f_{ij}) + e_n \quad (6)$$

Where,  $R_n$  is the reflectance for the  $i^{\text{th}}$  pixel,  $r_j$  is the spectral reflectance of the  $j^{\text{th}}$  surface component,  $f_{ij}$  is the fraction of the  $j^{\text{th}}$  surface component in the  $i^{\text{th}}$  pixel, and  $e_n$  denotes the pixel noise term. Following Smith *et al.* (*In Press*), generic example spectrums of green vegetation, non-photosynthetic vegetation (NPV), and char (as presented in Smith *et al.* 2005) were applied to the "ENVI linear spectral unmixing" tool as these example spectrums are

broadly similar across most vegetation types (Elvidge 1990, Landmann 2003). The same generic endmembers were used for all of the fires.

Pixel values from the five burn severity index images and three fractional cover images were extracted at the field subplot center point locations using an ArcInfo GRID AML (coded by J. Evans). For simplicity throughout the remainder of this paper, we may refer to all eight image variables as indices, although strictly speaking fractional cover images are not indices but physically-based estimates.

### Field Sampling

Fire effects data were collected 11 Sep – 13 Oct, 2003 at the Missoula fires in Montana; 29 Sep – 23 Oct, 2003 at the Glacier fires in Montana; 6-14 Dec, 2003 in California; and 22-31 Jul, 2004 in Alaska. The goal for selecting field sites was to find burn severity conditions that were large enough to include many Landsat image pixels and be broadly representative of the range of post-fire conditions occurring across the post-fire landscape. When a desired burn severity condition that could be safely accessed was found, the center of the field site was placed a random distance away from and on a compass bearing perpendicular to the access road. While preliminary BARC maps were used as rough guides to navigate to burned areas of interest, field sites were classified as low, moderate, or high severity if the tree crowns were predominantly green, brown, or black, respectively. This field assessment of burn severity class often did not agree with the class shown on the preliminary BARC map. Each field site was intended to sample variation in fire effects *within* that burn severity condition, while other field sites were placed to sample variation in fire effects *between* different burn severity conditions.

Thus, the field site center was randomly located within a selected severity condition that was consistent in terms of observed fire effects and apparent pre-fire stand structure and composition. The site center was designated the center of plot A, while the remaining eight plots were laid out systematic distances away, with the site oriented according to slope direction (Fig. 2). The nine plots were intentionally spaced apart by unequal intervals of 20, 30, or 40 m to spread out the distribution of lag distances separating the observations. This was done to facilitate a more robust assessment of the spatial autocorrelation in fire effects that underlies observed patterns, which will be a topic for another paper. Each plot was further subdivided into fifteen 1 m x 1 m subplots arrayed in a 3 row x 5 column grid as depicted in Fig. 2. The centers of plots A-I were all geolocated with a Global Positioning System (GPS), using a GeoExplorer (Trimble Navigation Limited, Sunnyvale, CA; Trade names are included for the benefit of the reader and do not imply endorsement by the U.S. Department of Agriculture or the University of Idaho.) to log a minimum of 150 positions, which were later differentially corrected and averaged to decrease the uncertainty to <2 m. Horizontal distances between plot centers were measured using a laser rangefinder to correct for slope effects. The centers of the other subplots were laid out using a cloth tape for distance and a compass for bearing, and marked with reusable pin flags. Subplot centers were not geolocated with the GPS, but their geolocations were later calculated based on their known systematic distance and bearing from the measured plot centers.

(Figure 2 about here)

A suite of fire effects were measured at the subplot or plot scale at each field site (Table 2). At the subplot scale, surface cover fractions of green vegetation, rock, mineral soil, ash, litter (new and old), and any large organics were estimated ocularly, with the aid of a 1 m<sup>2</sup> square quadrat. Percent char of each cover component was also recorded. At the plot scale, a ruler was used to measure depth of new litter (deposited post fire), old litter, and duff. Duff moisture was measured with a duff moisture meter (Robichaud *et al.* 2004 but only recorded in Alaska. Water repellency of soils that were charred lightly (black), moderately (gray), deeply (orange), or uncharred was measured using both a mini-disk infiltrometer (Decagon Devices, Pullman, WA; Trade names are included for the benefit of the reader and do not imply endorsement by the U.S. Department of Agriculture or the University of Idaho.) and a water drop penetration test (DeBano 1981). The water drop penetration test has been used more, but the mini-disk infiltrometer is considered a superior measurement because it is volumetric while the water drop test is not (Lewis *et al.* 2006). Water repellency was measured in only 7 of the 15 subplots within each plot (Fig. 2), to reduce the sampling effort required. Water repellency was not measured at 19 of the 50 sites sampled because the soils were too wet after recent precipitation for reliable measurements. A convex spherical densiometer was used to measure canopy closure around the center subplot (subplot 8), facing the four cardinal directions. Topographic features were also recorded at every plot, along with a digital photograph for reference.

At every site, percent canopy cover of grasses, forbs, low shrubs (<1 m tall, or <1 cm basal diameter if only a charred stub remained), and tree seedlings was estimated in a 1/750 ha circular plot; tall shrubs (>1 m tall, or >1 cm basal diameter if only a charred stub remained) and tree saplings were tallied in a 1/100 ha plot; and trees and snags (>12 cm dbh) were inventoried in a 1/50 ha plot. These three fixed-radius vegetation plots were arranged concentrically at site center (Fig. 2). With the exception of canopy closure measured at every square plot, the other overstory and understory vegetation variables were only assessed in one of the circular plots centered over plot A of each site (Fig. 2), and therefore comprised only 1/9 the sampling effort as the measures made at the 9 square plots A-I (Table 2).

(Table 2 about here)

### Analysis

A number of variables in Table 2 were not measured directly in the field but later calculated in Excel. Total organic charred and uncharred cover fractions were derived by combining the charred and uncharred fractions of the old litter and other organic constituents (stumps, logs, etc.) estimated in the field. Total inorganic charred and uncharred cover was similarly calculated by summing the charred and uncharred mineral soil and rock fractions estimated in the field. Total green, charred, and uncharred cover fractions, constrained to sum to unity to represent the entire 1 m x 1 m subplot, included all constituents except new litter (fallen since the time of burning). Water repellency measurements were weighted by the light, moderate, deep, and uncharred soil cover fractions they represented within the subplot sampled, then aggregated. The four canopy closure measurements made at each field plot with a convex spherical densiometer were rescaled from 0-96 canopy counts/measurement to 0-100%, as is standard with this instrument, and then averaged. Moss, liverwort, fern, forb, and low shrub cover estimates from the 1/750 ha center vegetation plot were summed to estimate total green (living), brown

(scorched), or black (charred) understory cover. The number of dead seedling, sapling, high shrub, and tree stems was divided by the total number of living and dead stems tallied in the field to calculate percent mortality. Field estimates of the green, scorched, or charred proportions of individual tree crowns were weighted by their crown length before averaging to represent the entire plot.

All of the eight image and 32 field variables were aggregated to the plot scale for this analysis. This was deemed necessary because the 2-m sampling interval between subplots was smaller than the positional uncertainty of the subplot locations. Also, not all variables were measured at the subplot scale, but all variables were measured at the plot scale (although not necessarily at every plot), and every plot position was geolocated with the GPS, making the plot scale most appropriate for this analysis of image-field data relationships. Correlation matrices between the image and field variables were generated in R (R Development Core Team, 2004), as were the boxplots and scatterplots that efficiently illustrate the highly variable relationships between the image and field variables.

## **Results**

### **Rapid Response**

Ours was a rapid response project, so it is useful and fair to ask how quickly image and field data were acquired after the fire. Each wildfire we sampled was wholly contained within a single Landsat TM Path (Table 1), which simplified the image acquisition variable to a single date. Thus on every wildfire, the distribution of days elapsed between burn date and image acquisition date, or field characterization date and image acquisition date, was wholly a function of the fire progression or the speed at which we characterized the plots in the field, respectively (Fig. 3). The California wildfires spread rapidly, driven by Santa Ana winds in predominantly chaparral vegetation. Fire progressions were typically slower in forests in Montana and Alaska, especially in Alaska where smoldering combustion in the deep duff can continue for months (Roger Ottmar, personal communication). The delay between when the field plots burned and when a cloud-free Landsat TM image was acquired ranged from 0-97 days (Fig. 3). In the case of the California and Montana fires (except Cooney Ridge), BAER teams needed BARC maps before a suitable Landsat TM image was available, prompting RSAC to obtain images from alternative sensors.

(Figure 3 about here)

The delay between when the plots burned and when they were characterized in the field ranged from 5-93 days (Fig. 3). The longest delays were in sampling the Robert and Wedge Canyon fires, because we were unable to access these fires west of Glacier National Park until after we had finished sampling the Black Mountain and Cooney Ridge fires near Missoula. The Missoula, Glacier, California, and Alaska wildfires are arranged top-bottom along the y-axes in Fig. 3 by the order in which they were sampled. Generally, the speed of field work increased as the crews grew accustomed to the sampling protocol.

The dates of image acquisition and field plot characterization were fairly well balanced (Fig. 3). The field plots at the Cooney Ridge and two California wildfires were imaged 43-19 days before we could reach them on the ground, while field plot sampling at the other five wildfires followed image acquisitions by 1-48 days (Fig. 3).

We tested if any of the time lags between burning, image acquisition, and field characterization dates might confound the image-field data relationships of primary interest. Plot-level correlation matrices were generated between the eight image variables and the 14 surface cover fractions measured or calculated at all of the subplots. Of the 418 field plots sampled in this study, 117 had no variation in 30 m pixel values between the subplots, preventing calculation of correlation matrices within these. The mean of the correlations (absolute values) for each of the remaining 301 plots was plotted against the three time lag variables to form three scatterplots, and smoothed loess functions were fit to these scatterplots to illustrate the trends (Fig. 4). These trends were tested for significance, but none proved to be significant over the ranges of time lags shown.

(Figure 4 about here)

### Image Indices

The distributions of correlations between the eight image-derived variables and 32 fire effects measures from the field varied somewhat between the eight wildfires sampled, but were usually not statistically significant when compared across the eight image variables (Fig. 5). When the correlation matrices from the Montana, California, and Alaska regions were combined, so that the three regions were apportioned equal weight despite their different plot counts, it became more evident that green fraction and NBR performed best across these three very different ecosystems (Fig. 6). Green fraction was more highly correlated to NBR than any of the other image indices in Montana ( $r = 0.82$ ), California ( $r = 0.63$ ), and Alaska ( $r = 0.86$ ).

(Figure 5 about here)

(Figure 6 about here)

### Fire Effects

The distributions of correlations between the 32 fire effects measures and the eight image variables were highly variable when compared across the fire effects variables (Figs. 7-9). No fire effects were consistently highly correlated to any of the image variables across all fires. The overstory measures of canopy closure, green and charred tree crowns were most highly correlated to the image variables in Montana, while understory measures had less influence (Fig. 7). Old litter depth was more highly correlated to the image variables than other surface or subsurface measures, which varied widely in correlation strength (Fig. 7).

(Figure 7 about here)

The poorest correlations generally among the three regions sampled were observed in California (Fig. 8). Trees were usually lacking in this predominantly chaparral vegetation, making tall shrub

and sapling mortality the best vegetation correlates to the image variables. Water repellency as measured by infiltrometer rate was better correlated to the image variables here than in Montana or Alaska (Figs. 7-9), perhaps because the relative lack of vegetation and litter/duff cover in California exposed so much more soil here than in the other regions. However, spectral differentiation between the organic and inorganic surface components was poor (Fig. 8).

(Figure 8 about here)

In Alaska, canopy closure and green tree crown overstory measures, and understory tall shrubs, were more highly correlated to the image variables than were other vegetation effects (Fig. 9). Charred and uncharred organics were correlated relatively well, as were total green, charred, and uncharred cover fractions, because little inorganic fraction existed to confuse the spectral reflectance signal in Alaska compared to Montana and California (Figs 7-9). The deep surface organic layer appears to be a more influential driver of spectral reflectance here in the black spruce forests of Alaska than in the ecosystems we sampled in Montana or California.

(Figure 9 about here)

We present more detailed results for exposed mineral soil cover, a measure of particular interest to BAER teams, to exemplify the high variability in observed fire effects between our selected study regions. Percent soil cover was twice as prevalent in California (68 %) than in Montana (33 %), while it comprised <9% of total surface cover in Alaska (Table 3). Although not shown here, surface organics followed the opposite trend. Aggregating the soil subplot measures to the plot and site levels improves the calculated correlations to the image variables by better capturing both the fine-scale (subpixel) variability in soil cover sampled on the ground, as well as the moderate-scale variability between pixels within the sampled burn severity condition (Fig. 2). However, aggregation reduces the sample size (N) at an even faster rate, hence producing fewer significant correlations at the plot scale, and fewest at the site scale (Table 3).

(Table 3 about here)

## **Discussion**

### **Rapid Response**

*Satellite Measures.* Landsat images were often not immediately available when BAER teams on the ground had critical need for BARC maps. SPOT imagery was used for BARC maps at the Montana fires. Coarse-resolution MODIS imagery was first used at the California fires, which was later supplanted by a higher-resolution MASTER (airborne) image at the Simi fire, and an ASTER image at the Old fire (Clark et al. 2003).

Further discussion of alternative satellite sensors to Landsat is warranted given that other sensors are used, and because the aging Landsat 5 and 7 satellites will inevitably fail like their predecessors, which is likely to precede the successful launch of a replacement from the Landsat Data Continuity Mission (<http://ldcm.nasa.gov/>). The SPOT 4 and 5 satellites have a short-wave infrared (SWIR) band 4 (1580-1750 nm) that closely approximates Landsat SWIR band 5 (1550-

1750 nm), rather than the Landsat SWIR band 7 (2080-2350 nm) preferred for calculating NBR (Eq. 1). In theory, the SPOT SWIR band 4 may be less useful for burn severity mapping using NBR than Landsat band 7 (Key and Benson 2003b). At the Cooney Ridge fire, we compared NBR calculated from the 31 Aug 2003 Landsat 5 image used in this paper to NBR calculated from a 1 Sep 2003 SPOT 4 image acquired the next day. A paired t-test showed that Landsat 5-derived NBR did only insignificantly better ( $p = 0.25$ ) than SPOT 4-derived NBR in terms of mean correlation strength to our 32 fire effects measures, although Landsat 5 NDVI did significantly better ( $p < 0.0001$ ) than SPOT 4 NDVI, perhaps because the latter image appeared more smoky. We repeated these statistical tests at the Old fire, where a 18 Nov 2003 ASTER image was acquired one day prior to the 19 Nov 2003 Landsat 5 image used in this analysis. In this case, the ASTER SWIR band 6 (2185-2225 nm) very closely approximates Landsat SWIR band 7. Here, the mean correlation strength of Landsat NBR to the 32 field measures was only negligibly higher ( $p = 0.73$ ) than with ASTER NBR, while ASTER NDVI did only negligibly better than Landsat NDVI ( $p = 0.95$ ). Finally in Alaska, we took advantage of a 9 Sep 2004 Landsat 7 ETM+ image acquired one day after the 8 Sep 2004 Landsat 5 TM image to compare these two sensors. Fortunately, all of our field plots happened to be situated along the center of the ETM+ scene, the portion of the satellite path unaffected by the data gaps in ETM+ imagery since the 31 May 2003 failure of the scan line corrector (<http://landsat7.usgs.gov/updates.php>). We found Landsat 7 did significantly better than Landsat 5 using NBR ( $p = 0.002$ ), but only insignificantly better using NDVI ( $p = 0.89$ ). Using a 3 Aug 2002 pre-fire ETM+ scene to also calculate delta indices, we also found significant improvements using Landsat 7 dNBR ( $p < 0.0001$ ) and RdNBR ( $p = 0.0008$ ), but only insignificant improvement using Landsat 7 dNDVI ( $p = 0.30$ ). We attribute the better Landsat 7 results to the improved radiometric resolution and less degraded condition of the ETM+ sensor compared to Landsat 5 TM.

*Field Measures.* Fire effects can change quickly. Ash cover in particular is rapidly redistributed by wind and rainfall, which may largely explain why it was a consistently poor correlate to the image variables (Figs. 7-9). Needlecast (new litter) on moderate severity sites is another dynamic phenomenon, as is green vegetation regrowth. In only two weeks post fire, bear grass appeared in Montana; manzanita, chamise, and other chaparral shrubs sprouted in California; and fireweed sprouted in Alaska and Montana. We made every effort to characterize our field sites as quickly after burning as was safely possible, but at a rate of one field site characterized per day, plus travel time just to get to the fire, it proved difficult to complete the fieldwork at each fire before weather and vegetation recovery had altered the post-fire scene. Indeed, the plot-level correlation strength between image and field data diminished as more days elapsed between field plot burning and field plot characterization (Fig. 4). Although this proved insignificant when assessed out to the maximum delay of 93 days ( $p = 0.35$ ), the 0-32 day portion of the trend was significant based on 91 field plots ( $p = 0.008$ ). Thus conditions on the ground may change more rapidly for about the first month after a fire than afterwards. This addressed a concern expressed by Hudak *et al.* (2004b) over fire effects changing before they can be characterized in the field, but the weakness of this trend does more to alleviate this concern since it is apparently not an overriding factor.

### Burn Severity Indices

Several other indices were tested in a preliminary analysis of this same dataset (Hudak *et al.* 2006) but were excluded from the results presented in this paper because they were unhelpful. Hudak *et al.* (2006) found that neither the Enhanced Vegetation Index (EVI) (Huete *et al.* 2002) nor delta EVI performed significantly better than the NBR or dNBR, respectively. Hudak *et al.* (2006) also tested three mathematical manipulations of the NBR that include the Landsat thermal band 6 (Holden *et al.* 2005), along with their respective delta indices, but these also performed no better than simple NBR and dNBR (Hudak *et al.* 2006).

The NBR index currently used by RSAC and EROS appears to be slightly more satisfactory than NDVI, not only because it performed slightly better in our broad analysis, but because the SWIR band used in the NBR formula (Eq. 1) is less vulnerable to the scattering effects of smoke and haze particulates than the RED band used to calculate NDVI (Eq. 4). On the other hand, NDVI did not perform significantly worse than NBR, meaning NDVI could be substituted for NBR if, for example, Landsat or other imagery with a SWIR band were unavailable.

Hudak *et al.* (2006) found the NBR and NDVI indices outperformed their corresponding delta indices, as well as RdNBR, at the Montana and Alaska fires. On the other hand, the delta indices performed better at the California fires. We analyzed the dataset in greater detail here by partitioning the results by individual fires (Fig. 5), but the same general conclusions still hold. The RdNBR performed best at the Old fire in California, which among the vegetation types we sampled was probably most similar structurally to the vegetation in the Sierra Nevada Mountains where RdNBR was successfully applied (Miller and Thode *In Press*). However, few differences between burn severity indices were significant at any of the fires (Fig. 3), which was our rationale for combining the correlation matrices from the Montana, California, and Alaska regions to assess which indices best correlated to fire effects over these very different ecosystems (Fig. 6). NBR produced the highest correlations overall, although by a negligible margin over NDVI. RSAC and EROS use the dNBR by default to produce BARC maps, which should be advantageous over NBR for extended (instead of immediate) burn severity assessment (Key and Benson 2003b). The influence of immediate fire effects on pixel reflectance attenuates over time, so including a pre-fire image to map the magnitude of change in the scene is advantageous over simply mapping post-fire condition (Hudak *et al.* 2007). RSAC and EROS currently archive the continuous pre-fire NBR, post-fire NBR, and dNBR products, and our results support this as a sensible practice.

### Fire Effects

Green vegetation had more influence on the image variables than any other surface constituent (Figs. 5-9). In other words, burn severity as indicated in BARC maps is more sensitive to vegetation effects than soil effects, so we caution against interpreting BARC maps as burn severity maps, particularly in terms of soil severity (Parsons and Orlemann 2002), without the field verification that BAER teams often do to improve their interpretation of burn severity from these maps. It should be expected that indices derived from a satellite image overhead will correlate better with vegetation than with soil characteristics, because vegetation occludes the ground. Percent soil cover was most highly correlated to NBR in Montana and to dNDVI in California and Alaska (Table 3). The strength of these correlations corresponded with the degree of soil exposure: strongest in California, less strong in Montana, and weakest in Alaska.

Similarly, surface organics that protect the soil were more prevalent in Alaska than in Montana or especially California, again matching the trend in their correlation strength to the image variables (Figs. 7-9). We also found burn severity indices generally correlated better to surface variables than soil water repellency variables (Figs. 7-9), since surface reflectance should be more correlated to surface characteristics than to soil processes measured beneath the surface.

Hudak *et al.* (2004a) used semivariograms to show that fire effects can vary greatly across multiple spatial scales ranging from the 2-m sampling interval between adjacent subplots to the 130-m span of an entire field site (Fig. 2). Our study supports earlier evidence that fire effects are more heterogeneous on low and moderate severity sites than on high severity sites (Turner *et al.* 1999). This helps explain why higher severity sites are more accurately classified on BARC maps than lower severity sites (Bobbe *et al.* 2003). Observed pixel reflectance is a function of all of the constituents occurring within that pixel, and to a lesser degree its neighboring pixels. Thus an image index of burn severity should not be expected to correlate very well with any single fire effect measured on the ground, when pixel reflectance is a function of multiple fire effects.

Pixels are fundamentally mixed, which is our primary justification for pursuing spectral mixture analysis (SMA) as a more robust strategy for mapping fire effects. The green, brown, and black variables included with the five burn severity indices in our analysis (Figs. 5-6) represent green vegetation, dead or non-photosynthetic vegetation (NPV), and char fractional cover, respectively, estimated via SMA. The six multispectral bands of Landsat TM images are not nearly as suited for SMA as hyperspectral images, but sufficient for estimating green and possibly char fractions. Our green fractional cover estimate performed negligibly better than NBR (Figs. 5-6), the best of the burn severity indices tested. Perhaps most importantly, an image of estimated green vegetation cover has direct biophysical meaning, even to a non-expert, unlike NBR or any other index. Importantly, calculation of the green (and other) fractions by spectral mixture analysis does not rely on the inclusion of all six Landsat bands. In the event that imagery is only available with the spectral equivalent of say, Landsat bands 1-4, these cover fractions could still be produced, whereas the NBR could not. Our inclusion of the green, brown, and black fractional cover estimates is instructive as it suggests that the five burn severity indices tested correspond more closely to green and black fractions than to brown fraction, the poorest correlate to measured fire effects. Unlike ash cover, char fraction remains relatively intact in the post-fire scene, and might be a suitable biophysical variable upon which to base a burn severity map derived using SMA. Green vegetation fraction is even more highly correlated with fire effects (Figs. 5-6) and NBR than char fraction, so a map of green vegetation fraction could be just as useful to BAER teams as the current dNBR-based BARC maps. For instance, it would be more difficult to misconstrue a green cover fraction map (labeled as such) as a soil severity map, as NBR- or dNBR-based BARC maps have been misinterpreted (Parsons and Orlemann 2002). An added advantage of mapping fractional cover estimates is that they represent remote analogues to the traditional field 'severity' measures of percent green, brown, and black (Lentile *et al.* 2006).

An important consideration not conveyed in Figs. 7-9 is the tremendous variation in inorganic and organic surface cover fractions observed in this study both within and between study regions. Lentile *et al.* (this issue) details the variability in ash, soil, and surface organics in each region and across low, moderate, and high severity burn classes. Ash cover was highest in Montana, soil cover in California, and surface organics in Alaska. The broader reach of the

boxplots in Figs. 7-8 compared to Fig. 9 is in large part due to greater variability in inorganic and organic cover fractions in Montana and California than in Alaska, where much of the organic matter persists even after a severe burn. Conversely, little soil was exposed in Alaska compared to Montana or especially California. We tested whether correlations between percent soil and the image variables were higher on high severity sites. Generally they were not, except in Alaska, because so little soil was exposed except on high severity sites (Lentile *et al.* this issue). Summarizing all of the fire effects measured in the field is not essential for communicating the image-field data relationships central to this paper, yet it is helpful to remember that sampling a wide range of variability in fire effects in the field is essential for understanding the influence any of them may have on pixel reflectance, because of scale effects related to the spectral mixing issue already described.

We chose percent exposed mineral soil, probably the most important fire effects measure to BAER teams (H. Shovic, *personal communication*), to further illustrate this scaling issue. When percent soil is correlated to pixel values at the fundamental subplot scale, often only 1 or at most 4 Landsat TM pixels will be sampled in a 9 m x 9 m plot. Aggregating the data from the 15 subplots to the plot scale invariably improves the correlations (Table 3), with the multiple subplot locations providing appropriate weight to the pixels being aggregated. Further aggregation to the site scale further improves the correlations because that much more pixel variation gets sampled (Table 3). This same pattern can be observed for the other fire effects sampled under our spatially nested design (Fig. 2). Many validation studies may find poor relationships because widely separated points only sample single pixels in the image product subject to validation, and single points are not very representative of image pixel values if the field variables of interest vary greatly at subpixel scales.

The downside of nested sampling designs such as ours is fewer field sites get sampled across the landscape, because so much time is required to characterize a single site. We did not aggregate our data to the site scale for the bulk of our analysis (Figs. 5-9) because that would have left too few sample units upon which to base the correlations, especially at the level of individual wildfires (Fig. 5). Table 3 uses percent soil cover to exemplify the scale considerations that are important for statistically analyzing nested sample data. The significance tests at the subplot scale have unreliably inflated degrees of freedom stemming from spatial autocorrelation, because the same image pixels get oversampled by the closely spaced subplots. On the other hand, the significance tests at the plot scale are reliable because plots A-I are spaced far enough apart to ensure that different image pixels get sampled (Fig. 2), although this does not completely eliminate potential autocorrelation effects (Hudak *et al.* 2004a). In summary, our hierarchical sampling strategy is not recommended for BAER teams or others who require rapid response or rapid results. It was designed specifically to explore and exploit the spatial autocorrelation that underlies observed patterns in fire effects, a topic for a future paper.

While a soil fractional cover map would be very useful to BAER teams, spectra for most soil types would be well mixed with the spectra for surface organics and NPV, and difficult to differentiate with 6-band multispectral Landsat data. Soil types varied greatly within and between our sample sites, making application of a single, generic soil endmember untenable. A spectral library of endmember spectra for specific soil types is available online from USGS (<http://pubs.usgs.gov/of/2003/ofr-03-395/datatable.html>). We used a field spectroradiometer

(ASD FieldSpecPro FR) to gather endmember spectra of soils, char, ash, NPV, litter, and the major plant species at all of our sampled wildfires, and will make these available in an online spectral library much like the USGS has. We also obtained airborne hyperspectral imagery over all of our field sites in the eight wildfires sampled, to test more rigorously whether maps of green, char, soil, or other cover fractions derived from spectral mixture analysis could potentially replace current burn severity maps based on indices (Lewis *et al.* this issue). Moreover, the 4-5 m spatial resolution of our hyperspectral imagery will allow more rigorous assessment of the tremendous spatial heterogeneity occurring at subpixel scales (i.e., 30 m Landsat). Our spatially nested field data will serve as valuable ground truth for validating estimates of multiple cover fractions derived from spectral mixture analysis.

## Conclusions

Our results show that the NBR and dNBR burn severity indices, as are currently used in BARC maps of large wildfires in the United States, are sound choices for rapid, preliminary assessment of immediate post-fire burn severity across different ecosystems. We recommend that RSAC and EROS continue their current practice of archiving the continuous NBR and dNBR layers upon which BARC maps are based, for future retrospective studies. The correlations of NDVI and dNDVI to the same suite of 32 fire effects measures generally were not significantly worse, so these indices could serve as suitable substitutes for NBR and dNBR. The Monitoring Trends in Burn Severity project should consider extending their historical scope and apply NDVI or dNDVI to map burn severity on fires preceding the availability of Landsat TM imagery, or from the 1972-84 Landsat Multispectral Scanner (MSS) image record, since MSS images lack the SWIR band needed to calculate NBR. The RdNBR only produced better correlations to fire effects at one of the eight wildfires sampled; thus, it may have more limited broad-scale utility. Time and money limited our sampling to only three regions, but it is difficult to imagine three ecosystems that could be more representative of the diverse fire ecology in North America than the three we selected in western Montana, southern California, and interior Alaska.

Our results show high variability in the relationships between indices derived from satellite imagery and fire effects measured on the ground, yet some remarkable consistencies across the three ecosystems sampled. First, none of the indices were very highly correlated with any of the 32 specific fire effects measured, which is likely a reflection of the 30-m scale of Landsat data relative to the finer scale at which fire effects vary. Second, the uppermost vegetation layers (trees in Montana and Alaska, shrubs in California) had more influence on the image indices than surface vegetation. Likewise, larger surface cover fractions were better correlated to the image indices than minor surface cover fractions or subsurface measures of soil water repellency. BAER teams should consider BARC maps much more indicative of post-fire vegetation condition than soil condition, and factor in that awareness when validating BARC maps on the ground. In terms of soil effects, BARC maps are more likely to be inaccurate on low or moderate severity sites than on high severity sites, where less vegetation remains to obstruct the view of the soil condition from above, as in an image.

By definition, burn severity is a measure of the ecological changes wrought by fire. Understory vegetation response (Lentile *et al.* this issue, and references therein), potential for soil erosion, and probable effects on soil nutrients are all more pronounced where fires have consumed more

fuel and resulted in more mineral soil exposure. Our results demonstrate that image spectral reflectance upon which indices are based is largely a function of the proportions of surface materials comprising the scene. Our future research will focus on spectral mixture analysis of hyperspectral imagery. We think this approach acknowledges that pixel reflectance is fundamentally a mixture of charred, dead, green, and nonphotosynthetic vegetation, soil, rock and ash materials that are highly variable at fine scales.

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